# Recent results for the 3D Quasi-Geostrophic Equation 

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## Physical model

- The Quasi-Geostrophic system of equations models the evolution of the temperature in the atmosphere.
- It can be rigorously derived from the Primitive Equations (Euler equation with Coriolis force and Boussinesq approximation, see Bourgeois Beale (94) and Desjardins Grenier 98)
- At large scale, this Rossby effect is very important. Asymptotically, this leads to the so-called geostrophic balance which enforces the wind velocity to be orthogonal to the gradient of the pressure in the atmosphere (see Pedlosky).
- This model is extensively used in computations of oceanic and atmospheric circulation, for instance, to simulate global warming.


## The unknown and parameters

- The dynamic is encoded in $\Psi$, the stream function for the geostrophic flow.
- That is, the 3D velocity $(w, U)=(0, u, v)$ has its horizontal component verifying

$$
(u, v)=\left(-\partial_{x_{2}} \Psi, \partial_{x_{1}} \Psi\right), \text { or in short }: U=\bar{\nabla}^{\perp} \Psi
$$

where we denote

$$
\bar{\nabla} \Psi=\left(0, \partial_{x_{1}} \Psi, \partial_{x_{2}} \Psi\right)
$$

- From the model, the buoyancy is given by

$$
\Theta=\partial_{z} \psi
$$

- We denote

$$
\nabla_{\lambda} \phi=\left(\lambda \partial_{z} \phi, \partial_{x_{1}} \phi, \partial_{x_{2}} \phi\right), \quad L_{\lambda} \phi=\operatorname{div}\left(\nabla_{\lambda} \phi\right)
$$

where $\lambda=-1 / \Theta_{z}^{0}$, is a given function, of $z$ only, associated to the buoyancy of a reference state.

## The equation

The function $\Psi$ is solution to the following Initial Boundary value problem:

$$
\begin{aligned}
& \left(\partial_{t}+\bar{\nabla}^{\perp} \psi \cdot \nabla\right)\left(L_{\lambda} \Psi+\beta_{0} x_{2}\right)=0, \quad t>0, \quad z>0, \quad x \in \mathbb{R}^{2}, \\
& \left(\partial_{t}+\bar{\nabla}^{\perp} \psi \cdot \nabla\right) \gamma_{\nu}\left(\nabla_{\lambda} \Psi\right)=\nu \bar{\Delta} \psi, \quad t>0, \quad z=0, \quad x \in \mathbb{R}^{2} \\
& \Psi(0, z, x)=\Psi^{0}(z, x) . \quad t=0, \quad z>0, \quad x \in \mathbb{R}^{2} .
\end{aligned}
$$

The parameter $\beta_{0}$ comes from the usual $\beta$-plane approximation. The term $\gamma_{\nu}\left(\nabla_{\lambda} \Psi\right)$ stands for the Neumann condition at $z=0$ associated to the operator $L_{\lambda} \Psi$. If $\lambda$ is regular, this coincides with $-\lambda(0) \partial_{z} \psi(0, \cdot)$. The $\nu$ term is due to the Ekman pumping. $\nu=0$ corresponds to the inviscid case.

- Both, the value of the elliptic operator $L_{\lambda} \Psi$, and the Neumann condition $\gamma_{\nu}\left(\nabla_{\lambda} \Psi\right)$ at the boundary $z=0$, are advected by the stratified flow with velocity $U=\bar{\nabla}^{\perp} \Psi$. At each time, $\Psi$ can be recovered, solving the boundary value elliptic equation.
- Main difficulty: treatment of the boundary condition.


## The inviscid case

We assume that $\nu=0$, and that there exists $\Lambda>0$ such that

$$
\frac{1}{\Lambda} \leq \lambda(z) \leq \Lambda, \quad \text { for } z \in \mathbb{R}^{+}
$$

## Theorem (Puel-V.)

Consider an initial value $\Psi^{0}$ such that

$$
L_{\lambda} \Psi^{0} \text { and } \nabla_{\lambda} \psi^{0} \text { are in } L^{2}\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right), \quad \gamma_{\nu}\left(\nabla_{\lambda} \Psi^{0}\right) \in L^{2}\left(\mathbb{R}^{2}\right) .
$$

Then, there exists $\Psi$ weak solution to the Quasi-Geostrophic equation on $(0, \infty) \times \mathbb{R}^{+} \times \mathbb{R}^{2}$, such that for every $T>0$, $\nabla_{\lambda} \Psi \in L^{\infty}\left(0, T ; L^{2}\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)\right) \cap C^{0}\left(0, T ; L_{\text {loc }}^{2}\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)\right)$.

Novack recently extended the theory to general $L^{p}$.

## Lateral boundary conditions

We consider a domain of the form $\mathbb{R}^{+} \times \Omega$, where $\Omega \subset \mathbb{R}^{2}$ is bounded. Think, for instance about a rotating box filled with a fluid.

## Theorem (Novack-V.)

The natural lateral boundary conditions on $\mathbb{R}^{+} \times \partial \Omega$ are

$$
\begin{aligned}
& \Psi \text { depends only on } z \text { on } \mathbb{R}^{+} \times \partial \Omega, \\
& \frac{d}{d t} \int_{\partial \Omega} \partial_{\nu} \Psi d \hat{x}=0 .
\end{aligned}
$$

We can also construct global weak solutions of $Q G$ with the addition of these boundary conditions.

This corresponds to a partial Dirichlet condition (up to the dependency on z), together with a mean value of Neumann condition on $\partial \Omega$.

## The case with Ekman pumping

We assume that $\lambda(z)=1$, and $\nu>0$.

## Theorem (Novack-V.)

Consider an initial value $\nabla \psi^{0} \in L^{2}\left(\mathbb{R}_{+}^{3}\right) \cap H^{p}\left((0, \infty) \times \mathbb{R}^{2}\right)$ with $p \geq 3$.
Then, there exists a unique global solution $\nabla \psi$ to the Quasigeostrophic equation on $(0, \infty) \times \mathbb{R}^{+} \times \mathbb{R}^{2}$, such that for every $T>0, \nabla_{\lambda} \Psi \in C^{0}\left(0, T ; H^{p}\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)\right)$.

Especially, if the initial is smooth $\left(C^{\infty}\right)$, then the unique solution is also smooth.

## Main difficulty

To simplify the exposition, let us consider the case with out forcing with $\beta=0$, and $\lambda=1$.

$$
\begin{aligned}
& \left(\partial_{t}+\bar{\nabla}^{\perp} \Psi \cdot \nabla\right)(\Delta \Psi)=0, \quad \text { for } z>0, \\
& \left(\partial_{t}+\bar{\nabla}^{\perp} \Psi \cdot \nabla\right)\left(\partial_{z} \Psi\right)=0, \quad \text { for } z=0, \\
& \Psi(0, z, x)=\Psi^{0}(z, x) . \quad t=0 .
\end{aligned}
$$

- A priori estimates: for any $1 \leq p \leq \infty$ :

$$
\begin{aligned}
& \|\Delta \Psi(t)\|_{L \rho\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)} \leq\|\Delta \Psi(0)\|_{L \rho_{\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)},}, \\
& \left\|\partial_{z} \Psi(t, 0)\right\|_{L \rho_{\left(\mathbb{R}^{2}\right)}} \leq\left\|\partial_{z} \Psi(0,0)\right\|_{L \rho\left(\times \mathbb{R}^{2}\right)},
\end{aligned}
$$

- No compactness on the trace of $\partial_{z} \psi$ at $z=0$ !


## A special case: the Surface Quasi Geostrophic Equation

- If $\Delta \Psi(0)=0$, then $\Delta \Psi(t)=0$ for all $t \geq 0$.
- Denote $\theta=\partial_{z} \Psi$ defined at $z=0$. Then $\theta$ is solution to

$$
\begin{align*}
& \partial_{t} \theta+U . \nabla \theta=0, \quad t>0,(x, y) \in \mathbb{R}^{2}  \tag{1}\\
& \theta=\theta_{0}, \quad t=0,(x, y) \in \mathbb{R}^{2} \tag{2}
\end{align*}
$$

and the velocity $U$ can be expressed in $\mathbb{R}^{2}$, via a nonlocal operator, as

$$
U=\nabla^{\perp} \Delta^{-1 / 2} \theta
$$

- This model has been popularized as a toy problem for 3D fluid mechanics (see Constantin, Majda, Held, Pierrehumbert, Garner, Swanson ...).
- Our theorem extends to QG the result of Tabak for SQG, using different techniques.


## A new formulation (1)

- The proof does NOT use (and does not show) compactness on the trace of $\partial_{z} \psi$ at $z=0$.
- It is based on a reformulation of the problem into a system of equations (without equation on the trace).
- The stability (and compactness) for this problem is pretty simple.
- We then have to show the equivalence between the two formulations.


## A new formulation (2)

- Consider the Hodge decomposition in $L^{2}\left(\mathbb{R}^{+} \times \mathbb{R}^{2}\right)$ :

$$
u=\nabla_{\lambda} \phi+\operatorname{curl} v=\mathbb{P}_{\lambda} u+\mathbb{P}_{\text {curl }} u
$$

with curlv $\cdot \nu=0$ at $z=0$.

- The QG problem can be reformulated as

$$
\partial_{t} \nabla_{\lambda} \Psi+\mathbb{P}_{\lambda}\left(\bar{\nabla} \Psi^{\perp} \cdot \bar{\nabla} \nabla_{\lambda} \Psi\right)=0, \quad \text { on } \mathbb{R}^{+} \times \mathbb{R}^{2} \times \mathbb{R}^{+} .
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$$

- Taking the div of the equation gives the first QG equation, thanks to

$$
\operatorname{div}\left(\mathbb{P}_{\lambda} \cdot\right)=\operatorname{div}(\cdot), \quad \partial_{i}(\bar{\nabla} \Psi)^{\perp} \cdot \bar{\nabla} \partial_{i} \Psi=0 .
$$

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$$

- Taking the trace of the system a $z=0$ gives (formally) the trace condition of QG, since formally, at $z=0$

$$
\mathbb{P}_{\lambda}(f) \cdot \nu=f \cdot \nu
$$

## Similarity with the Euler equation

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Note that we have $\mathbb{P}_{\text {curl }}\left(\nabla_{\lambda} \Psi\right)=0$.

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$$

Note that we have $\mathbb{P}_{\text {curl }}\left(\nabla_{\lambda} \Psi\right)=0$.

- Euler Equation:

$$
\begin{aligned}
& \quad \partial_{t} \text { curlv }+\mathbb{P}_{\text {curl }}[\text { curlv } \cdot \nabla \text { curlv }]=0, \quad(t, x, z) \in \mathbb{R}^{+} \times \mathbb{R}^{2} \times \mathbb{R}^{+} . \\
& \text {with } \mathbb{P}_{\lambda}(\text { curlv })=0 \text { (that is curlv } \cdot \nu=0 \text { at } z=0 \text { ). }
\end{aligned}
$$

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- QG equation:

$$
\partial_{t} \nabla_{\lambda} \Psi+\mathbb{P}_{\lambda}\left(\bar{\nabla} \Psi^{\perp} \cdot \bar{\nabla} \nabla_{\lambda} \Psi\right)=0, \quad \text { on } \mathbb{R}^{+} \times \mathbb{R}^{2} \times \mathbb{R}^{+}
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Note that we have $\mathbb{P}_{\text {curl }}\left(\nabla_{\lambda} \Psi\right)=0$.

- Euler Equation:

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$$

with $\mathbb{P}_{\lambda}$ (curlv) $=0$ (that is curlv $\cdot \nu=0$ at $z=0$ ).

- The first equation of QG is equivalent to the vorticity equation of Euler:
- QG:

$$
\partial_{t} \operatorname{div} \nabla_{\lambda} \Psi+\bar{\nabla} \Psi^{\perp} \cdot \bar{\nabla}\left(\operatorname{div} \nabla_{\lambda} \Psi\right)=0
$$

- Euler:

$$
\partial_{t} \text { curlcurl } v+\text { curlv } \cdot \nabla(\text { curlcurl } v)=0 .
$$

## Proof of the Theorem

- Compactness holds for the reformulated problem.

Note that $\mathbb{P}_{\lambda}$ commutes with $\bar{\nabla}$, and is continuous in $L^{p}$.

- The two formulation of QG are equivalent.


## A special case: the Surface Quasi Geostrophic Equation

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- Denote $\theta=\partial_{z} \Psi$ defined at $z=0$. Then $\theta$ is solution to

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\begin{align*}
& \partial_{t} \theta+U . \nabla \theta=\nu \bar{\Delta} \Psi, \quad t>0,(x, y) \in \mathbb{R}^{2}  \tag{3}\\
& \theta=\theta_{0}, \quad t=0,(x, y) \in \mathbb{R}^{2} \tag{4}
\end{align*}
$$

and the velocity $U$ and the Ekman pumping term $\nu \bar{\Delta} \psi$ can be expressed in $\mathbb{R}^{2}$, via a nonlocal operator, as

$$
U=\nabla^{\perp} \Delta^{-1 / 2} \theta, \quad \nu \bar{\Delta} \Psi=\nu \Delta^{1 / 2} \theta
$$

- The propagation of regularity for this equation has first been proved by Kiselev, Nazarov and Volberg. The global regularity of solutions with initial values in $L^{2}$ has been proved first by Caffarelli V. Several other proofs has been proposed by Kiselev and Volberg, and Constantin and Vicol.


## The 3D case

- In the 3D case, the equation in $z>0$ is hyperbolic. We can have only propagation of regularity.
- We need the propagation of almost Lipschitz norm (possible log Lipschitz).
- The regularization effects on the boundary are only $C^{\alpha}$.


## Sketch of the proof (1)

We decompose the solution $\Psi=\Psi_{1}+\Psi_{2}$ into two components as follows:

$$
\left\{\begin{array} { l } 
{ \Delta \psi _ { 1 } = 0 } \\
{ \partial _ { \nu } \psi _ { 1 } = \partial _ { \nu } \psi }
\end{array} \quad \left\{\begin{array}{l}
\Delta \Psi_{2}=\Delta \psi \\
\partial_{\nu} \psi_{2}=0 .
\end{array}\right.\right.
$$

- The bulk of the proof is centered around verifying a version of the Beale-Kato-Majda criterion.


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\end{array} \quad \left\{\begin{array}{l}
\Delta \Psi_{2}=\Delta \psi \\
\partial_{\nu} \psi_{2}=0 .
\end{array}\right.\right.
$$

- The bulk of the proof is centered around verifying a version of the Beale-Kato-Majda criterion.
- The equation on the boundary of $\theta=\partial_{\nu} \Psi_{1}$ is of the form

$$
\partial_{t} \theta+u \cdot \bar{\nabla} \theta+(-\bar{\Delta})^{\frac{1}{2}} \theta=f,
$$

with $f=\bar{\Delta} \Psi_{2}$.

- The natural a priori bound for $f$ is in $B_{\infty, \infty}^{0}$.
- Using De Giorgi techniques, we get $\theta$ bounded in $C^{\alpha}$.


## Sketch of the proof (2)

- Bootstrapping an increase of regularity on the $C^{\alpha}$ on the drift-diffusion equation on the boundary gives that $\partial_{\nu} \psi \in L^{\infty}\left(0, T ; B_{\infty, \infty}^{1}\right)$ on the boundary.
- Using that the flow is stratified, this gives the "almost Lipschitz" bound needed on the velocity in $z>0$ generated by the boundary.


## Remark on the lateral boundary conditions

- In the case of the inviscid SQG, defined on a Bounded domain $\Omega \subset \mathbb{R}^{2}$, we need to define the velocity $U$.
- Constantin and Nguyen (17) proposed to define it through the Operator $\bar{\Delta}_{D}^{-1 / 2}$ with Dirichlet boundary condition 0 on $\partial \Omega$ :

$$
U=\bar{\nabla}^{\perp} \bar{\Delta}_{D}^{-1 / 2} \theta
$$

- This corresponds to a Dirichlet condition $\Psi=0$ on $\mathbb{R}^{+} \times \partial \Omega$ for the 3D QG.
- This is not the boundary condition derived from the primitive equation.
- The corresponding boundary condition for SQG can be retrieved using the Extension Operator of Caffarelli-Silvestre.


## Thank you

## Thank You !!

